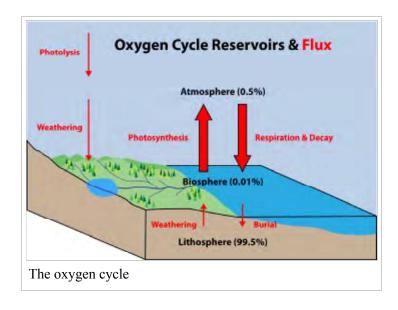
# Oxygen cycle

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The **oxygen cycle** is the biogeochemical cycle of oxygen within its three main reservoirs: the atmosphere (air), the total content of biological matter within the biosphere (the global sum of all ecosystems), and the Earth's crust. Failures in the oxygen cycle within the hydrosphere (the combined mass of water found on, under, and over the surface of planet Earth) can result in the development of hypoxic zones. The main driving factor of the oxygen cycle is photosynthesis, which is responsible for the modern Earth's atmosphere and life on earth (see the Great Oxygenation Event).

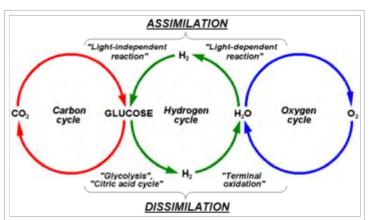


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## Reservoirs

By far the largest reservoir of Earth's oxygen is within the silicate and oxide minerals of the crust and mantle (99.5%). Only a small portion has been released as free oxygen to the biosphere (0.01%) and atmosphere (0.36%). The main source of atmospheric free oxygen is photosynthesis, which produces sugars and free oxygen from carbon dioxide and water:



Interconnection between carbon, hydrogen and oxygen cycle in metabolism of photosynthesizing plants

$$6 \text{ CO}_2 + 6 \text{H}_2 \text{O} + \text{energy} \longrightarrow \text{C}_6 \text{H}_{12} \text{O}_6 + 6 \text{ O}_2$$

Photosynthesizing organisms include the plant life of the land areas as well as the phytoplankton of the oceans. The tiny marine cyanobacterium Prochlorococcus was discovered in 1986 and accounts for more than half of the photosynthesis of the open ocean.<sup>[1]</sup>

An additional source of atmospheric free oxygen comes from photolysis, whereby high-energy ultraviolet radiation breaks down atmospheric water and nitrous oxide into component atoms. The free H and  $\overline{\mathbb{N}}$  atoms escape into space, leaving  $O_2$  in the atmosphere:

$$2 H_2O + energy \longrightarrow 4 H + O_2$$
  
 $2 N_2O + energy \longrightarrow 4 N + O_2$ 

The main way free oxygen is lost from the atmosphere is via respiration and decay, mechanisms in which animal life and bacteria consume oxygen and release carbon dioxide.

The lithosphere also consumes free oxygen by chemical weathering and surface reactions. An example of surface weathering chemistry is formation of iron oxides (rust):

$$4~{
m FeO} + {
m O}_2 \longrightarrow 2~{
m Fe}_2 {
m O}_3$$

Oxygen is also cycled between the biosphere and lithosphere. Marine organisms in the biosphere create calcium carbonate shell material (CaCO<sub>3</sub>) that is rich in oxygen. When the organism dies, its shell is deposited on the shallow sea floor and buried over time to create the limestone sedimentary rock of the lithosphere. Weathering processes initiated by organisms can also free oxygen from the lithosphere. Plants and animals extract nutrient minerals from rocks and release oxygen in the process.

## Capacities and fluxes

The following tables offer estimates of oxygen cycle reservoir capacities and fluxes. These numbers are based primarily on estimates from (Walker, J. C. G.<sup>[2]</sup>):

**Table 1**: Major reservoirs involved in the oxygen cycle

Reservoir	Capacity (kg O <sub>2</sub> )	Flux in/out (kg O <sub>2</sub> per year)	Residence time (years)
Atmosphere	$1.4\times10^{18}$	$3 \times 10^{14}$	4500
Biosphere	$1.6 \times 10^{16}$	$3 \times 10^{14}$	50
Lithosphere	$2.9 \times 10^{20}$	6 × 10 <sup>11</sup>	500 000 000

**Table 2**: Annual gain and loss of atmospheric oxygen (Units of  $10^{10}$  kg  $O_2$  per year)

Photosynthesis (land) Photosynthesis (ocean) Photolysis of N <sub>2</sub> O Photolysis of H <sub>2</sub> O	16,500 13,500 1.3 0.03			
Total gains				
Losses - respiration and decay				
Aerobic respiration Microbial oxidation Combustion of fossil fuel (anthropogenic) Photochemical oxidation Fixation of N <sub>2</sub> by lightning Fixation of N <sub>2</sub> by industry (anthropogenic) Oxidation of volcanic gases	23,000 5,100 1,200 600 12 10 5			
Losses - weathering				
Chemical weathering Surface reaction of O <sub>3</sub>	50 12			
Total losses	~ 30,000			

#### **Ozone**

The presence of atmospheric oxygen has led to the formation of ozone  $(O_3)$  and the ozone layer within the stratosphere:

$$\begin{array}{l} O_2 + uv \ light \longrightarrow 2 \ O \qquad (\lambda \lesssim 200 \ nm) \\ O + O_2 \longrightarrow O_3 \end{array}$$

The ozone layer is extremely important to modern life as it absorbs harmful ultraviolet radiation:

$${
m O_3 + uv \ light \longrightarrow O_2 + O} \hspace{0.5cm} (\lambda \lesssim 300 \ {
m nm})$$

### References

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- 2. Walker, J. C. G. (1980) The oxygen cycle in the natural environment and the biogeochemical cycles, Springer-Verlag, Berlin, Federal Republic of Germany (DEU).
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Categories: Ecology | Chemical oceanography | Photosynthesis | Biogeochemical cycle

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