# Seismology

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**Seismology** (pronunciation: /saiz mplədʒi/; from Ancient Greek  $\sigma$ εισμός (*seismós*) meaning "earthquake" and - $\lambda \alpha \gamma i \alpha$  (*-logía*) meaning "study of") is the scientific study of earthquakes and the propagation of elastic waves through the Earth or through other planet-like bodies. The field also includes studies of earthquake environmental effects such as tsunamis as well as diverse seismic sources such as volcanic, tectonic, oceanic, atmospheric, and artificial processes such as explosions. A related field that uses geology to infer information regarding past earthquakes is paleoseismology. A recording of earth motion as a function of time is called a seismogram. A seismologist is a scientist who does research in seismology.

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# History

Scholarly interest in earthquakes can be traced back to antiquity. Early speculations on the natural causes of earthquakes were included in the writings of Thales of Miletus (c. 585 BCE), Anaximenes of Miletus (c. 550 BCE), Aristotle (c. 340 BCE) and Zhang Heng (132 CE).

In 132 CE, Zhang Heng of China's Han dynasty designed the first known seismoscope.

In 1664, Athanasius Kircher argued that earthquakes were caused by the movement of fire within a system of channels inside the Earth.

In 1703, Martin Lister (1638 to 1712) and Nicolas Lemery (1645 to 1715) proposed that earthquakes were caused by chemical explosions within the earth.

The Lisbon earthquake of 1755, coinciding with the general flowering of science in Europe, set in motion intensified scientific attempts to understand the behaviour and causation of earthquakes. The earliest responses include work by John Bevis (1757) and John Michell (1761). Michell determined that earthquakes originate within the Earth and were waves of movement caused by "shifting masses of rock miles below the surface."<sup>[1]</sup>

From 1857, Robert Mallet laid the foundation of instrumental seismology and carried out seismological experiments using explosives. He is also responsible for coining the word "seismology".<sup>[2]</sup>

In 1897, Emil Wiechert's theoretical calculations led him to conclude that the Earth's interior consists of a mantle of silicates, surrounding a core of iron.<sup>[3]</sup>

In 1906 Richard Dixon Oldham identified the separate arrival of P-waves, S-waves and surface waves on seismograms and found the first clear evidence that the Earth has a central core.<sup>[4]</sup>

In 1910, after studying the 1906 San Francisco earthquake, Harry Fielding Reid put forward the "elastic rebound theory" which remains the foundation for modern tectonic studies. The development of this theory depended on the considerable progress of earlier independent streams of work on the behaviour of elastic materials and in mathematics.<sup>[5]</sup>

In 1926, Harold Jeffreys was the first to claim, based on his study of earthquake waves, that below the mantle, the core of the Earth is liquid.<sup>[6]</sup>

In 1937, Inge Lehmann determined that within the earth's liquid outer core there is a solid *inner* core.<sup>[7]</sup>

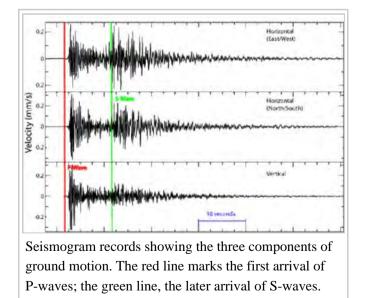
By the 1960s, earth science had developed to the point where a comprehensive theory of the causation of seismic events had come together in the now well-established theory of plate tectonics.

# Types of seismic wave

Seismic waves are elastic waves that propagate in solid or fluid materials. They can be divided into *body waves* that travel through the interior of the materials; *surface waves* that travel along surfaces or interfaces between materials; and *normal modes*, a form of standing wave.

### **Body waves**

There are two types of body waves, Pressure waves or Primary waves (P-waves) and Shear or Secondary waves (S-waves). P-waves are longitudinal waves that involve compression and expansion in the direction that the wave is moving and are always the first waves to appear on a seismogram as they are the fastest moving waves through solids. S-waves are transverse waves that move perpendicular to the direction of propagation. S-waves are slower than P-waves. Therefore, they



appear later than P-waves on a seismogram. Fluids cannot support perpendicular motion, so S-waves only travel in solids.<sup>[8]</sup>

#### **Surface waves**

The two main surface wave types are Rayleigh waves, which have some compressional motion, and Love waves, which do not. Rayleigh waves result from the interaction of vertically polarized P- and S-waves that satisfy the boundary conditions on the surface. Love waves can exist in the presence of a subsurface layer, and are only formed by horizontally polarized S-waves. Surface waves travel more slowly than P-waves and S-waves; however, because they are guided by the Earth's surface and their energy is thus trapped near the surface, they can be much stronger than body waves, and can be the largest signals on earthquake seismograms. Surface waves are strongly excited when their source is close to the surface, as in a shallow earthquake or a near surface explosion.<sup>[8]</sup>

#### Normal modes

Both body and surface waves are traveling waves; however, large earthquakes can also make the Earth "ring" like a bell. This ringing is a mixture of normal modes with discrete frequencies and periods of an hour or shorter. Motion caused by a large earthquake can be observed for up to a month after the event.<sup>[8]</sup> The first observations of normal modes were made in the 1960s as the advent of higher fidelity instruments coincided with two of the largest earthquakes of the 20th century - the 1960 Valdivia earthquake and the 1964 Alaska earthquake. Since then, the normal modes of the Earth have given us some of the strongest constraints on the deep structure of the Earth.

# Earthquakes

One of the first attempts at the scientific study of earthquakes followed the 1755 Lisbon earthquake. Other notable earthquakes that spurred major advancements in the science of seismology include the 1857 Basilicata earthquake, 1906 San Francisco earthquake, the 1964 Alaska earthquake, the 2004 Sumatra-Andaman earthquake, and the 2011 Great East Japan earthquake.

### **Controlled seismic sources**

Seismic waves produced by explosions or vibrating controlled sources are one of the primary methods of underground exploration in geophysics (in addition to many different electromagnetic methods such as induced polarization and magnetotellurics). Controlled-source seismology has been used to map salt domes, anticlines and other geologic traps in petroleum-bearing rocks, faults, rock types, and long-buried giant meteor craters. For example, the Chicxulub Crater, which was caused by an impact that has been implicated in the extinction of the dinosaurs, was localized to Central America by analyzing ejecta in the Cretaceous–Paleogene boundary, and then physically proven to exist using seismic maps from oil exploration.<sup>[9]</sup>

# **Detection of seismic waves**

Seismometers are sensors that detect and record the motion of the Earth arising from elastic waves. Seismometers may be deployed at the Earth's surface, in shallow vaults, in boreholes, or underwater. A complete instrument package that records seismic signals is called a seismograph. Networks of seismographs continuously record ground motions around the world to facilitate the monitoring and analysis of global earthquakes and other sources of seismic activity. Rapid location of earthquakes makes tsunami warnings possible because seismic waves travel considerably faster than tsunami waves. Seismometers also record signals from non-earthquake sources ranging from explosions (nuclear and chemical), to local noise from wind or anthropogenic activities, to incessant signals generated at the ocean floor and coasts induced by ocean waves (the global microseism), to cryospheric events associated with large icebergs and glaciers.

Above-ocean meteor strikes with energies as high as  $4.2 \times 10^{13}$  J (equivalent to that released by an explosion of ten kilotons of TNT) have been recorded by seismographs, as have a number of industrial accidents and terrorist bombs and events (a field of study referred to as forensic seismology). A major long-term motivation for the global seismographic monitoring has been for the detection and study of nuclear testing.

# Mapping the earth's interior

Because seismic waves commonly propagate efficiently as they interact with the internal structure of the Earth, they provide high-resolution noninvasive methods for studying the planet's interior. One of the earliest important discoveries (suggested by Richard Dixon Oldham in 1906 and definitively shown by Harold Jeffreys in 1926) was that the outer core of the earth is liquid. Since S-waves do not pass through liquids, the liquid core causes a "shadow" on the side of the planet opposite the earthquake where no direct S-waves are observed. In addition, P-waves travel much slower through the outer core than the mantle.

Processing readings from many seismometers using seismic tomography, seismologists have mapped the mantle of the earth to a resolution of several hundred kilometers. This has enabled scientists to identify convection cells and other large-scale features such as the large low-shear-velocity provinces near the core–mantle boundary.<sup>[10]</sup>

# Seismology and society

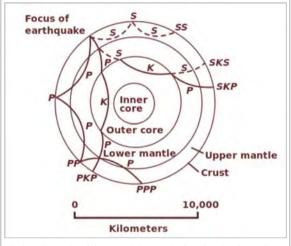
### **Earthquake prediction**

Forecasting a probable timing, location, magnitude and other important features of a forthcoming seismic event is called earthquake prediction. Various attempts have been made by seismologists and others to create effective systems for precise earthquake predictions, including the VAN method. Most seismologists do not believe that a system to provide timely warnings for individual earthquakes has yet been developed, and many believe that such a system would be unlikely to give useful warning of impending seismic events. However, more general forecasts routinely predict seismic hazard. Such forecasts estimate the probability of an earthquake of a particular size affecting a particular location within a particular time-span, and they are routinely used in earthquake engineering.

Public controversy over earthquake prediction erupted after Italian authorities indicted six seismologists and one government official for manslaughter in connection with a magnitude 6.3 earthquake in L'Aquila, Italy on April 5, 2009. The indictment has been widely perceived as an indictment for failing to predict the earthquake



Installation for a temporary seismic station, north Iceland highland.



Seismic velocities and boundaries in the interior of the Earth sampled by seismic waves and has drawn condemnation from the American Association for the Advancement of Science and the American Geophysical Union. The indictment claims that, at a special meeting in L'Aquila the week before the earthquake occurred, scientists and officials were more interested in pacifying the population than providing adequate information about earthquake risk and preparedness.<sup>[11]</sup>

#### **Engineering seismology**

Engineering seismology is the study and application of seismology for engineering purposes.<sup>[12]</sup> It generally applied to the branch of seismology that deals with the assessment of the seismic hazard of a site or region for the purposes of earthquake engineering. It is, therefore, a link between earth science and civil engineering.<sup>[13]</sup> There are two principal components of engineering seismology. Firstly, studying earthquake history (e.g. historical<sup>[13]</sup> and instrumental catalogs<sup>[14]</sup> of seismicity) and tectonics<sup>[15]</sup> to assess the earthquakes that could occur in a region and their characteristics and frequency of occurrence. Secondly, studying strong ground motions generated by earthquakes to assess the expected shaking from future earthquakes with similar characteristics. These strong ground motions could either be observations from accelerometers or seismometers or those simulated by computers using various techniques.<sup>[16]</sup>

# Tools

Seismological instruments can generate large amounts of data. Systems for processing such data include:

- CUSP (Caltech-USGS Seismic Processing)<sup>[17]</sup>
- RadExPro seismic software
- SeisComP3<sup>[18]</sup>

# Notable seismologists

- Aki, Keiiti
- Don L.
   Anderson
- Bolt, Bruce
- Claerbout, Jon
- Dziewonski, Adam Marian
- Ewing, Maurice
- Galitzine, Boris
   Borisovich

- Gamburtsev, Grigory A.
- Gutenberg,
  - Beno
- Hough, Susan
- Jeffreys,
- HaroldJones, Lucy
- Kanamori,
- Hiroo • Keilis-Borok,
- Vladimir ■ Knopoff,
  - Leon

- Lehmann,
- Inge Mallet,
- Robert ■ Mercalli,
- Giuseppe
- Milne, John
- Mohorovičić.
- Andrija
- Oldham, Richard
- DixonFusakichi Omori

- Sebastião de Melo,
   Marquia of
  - Marquis of Pombal
- Press, Frank
- Richards.
  - Paul G.
- Richter, Charles Francis
- Sekiya,
   Seikei
- Sieh, Kerry

- Paul G. Silver
- Tucker, Brian
- Vidale, John
- Wen,
- LianxingWinthrop,
- John
- Zhang Heng

- See also
  - Asteroseismology (starquakes)
  - Comparison of free geophysics software
- Cryoseism
- Engineering geology

- Forensic seismology
- Harmonic tremor
- Helioseismology
- IRIS Consortium
- Isoseismal map
- Linear seismic inversion
- Lunar seismology
- Quake (natural phenomenon), includes planets and stars

### Notes

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  /eost18389.pdf. External link in |title=(help)
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- 10. Wen & Helmberger 1998
- 11. Hall 2011
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- Seismic interferometry
- Seismic loading
- Seismic migration
- Seismic noise
- Seismic performance analysis
- Seismite
- Seismo-electromagnetics
- Seismotectonics
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# **External links**

 European-Mediterranean Seismological Center (http://www.emsc-csem.org/), real-time earthquake information website.



Wikimedia Commons has media related to *Seismology*.

- Seismological Society of America (http://www.seismosoc.org/).
- Incorporated Research Institutions for Seismology (http://www.iris.edu/).
- USGS Earthquake Hazards Program (http://earthquake.usgs.gov/earthquakes/).

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